

Mineralogical and Geochemical Appraisal of Clay Deposits in Papalanto and Its Environs, Southwestern, Nigeria

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Abstract

Three residual clay occurrences in Papalanto, Ifo, and Imoto areas which belong to the sedimentary basin of southwestern Nigeria were investigated to determine their industrial applications.

The samples were pulverized, sieved, digested with mineral acids and characterized. Clay mineralogy was determined using X-ray Diffraction (XRD). Elemental compositions of the clay samples were determined using Inductively Coupled Plasma-Mass Spectrometer (ICP-MS). Grain size distribution data were obtained by conducting grain size analysis in two parts; sieve analysis and sedimentation. Thermal properties, plasticity tests, density measurement, linear shrinkage and water absorption capacity were determined to evaluate their industrial potentials. Chemical Index of Alteration (CIA) was calculated to determine the degree of weathering in the area.

The X-ray diffraction results showed that kaolinite is the dominant mineral, while quartz, anatase and hematite are the major non clay minerals. Chemical data showed that the average values of SiO₂, Al₂O₃, and Fe₂O₃ were 59.46%, 22.16%, and 3.06% respectively constituting 98.3% of the bulk compositions. Papalanto possessed high plasticity and mouldability.

Evaluation of the clay thermal characteristics, firing colour, water absorption capacities and shrinkage values showed that the whitish Ifo clay and Papalanto kaolinitic clays could serve as raw materials for ceramics, building bricks, and other structural wares. Kaolin which is the dominant mineral in all the clays can be used for cosmetics, tooth paste, pharmaceutical purposes

Keywords: clay, structural wares, ceramics, anatase, building bricks

1. Introduction

Clay minerals typically form over long periods of time from the gradual chemical weathering of rocks, usually silicate-bearing, by low concentrations of carbonic acid and other diluted solvents. These solvents, usually acidic, migrate through the weathering rock after leaching through upper weathered layers. In addition to the weathering process, some clay minerals are formed through hydrothermal activity. There are two types of clay deposits: primary and secondary. Primary clays form as residual deposits in soil and remain at the site of formation and secondary clays are clays that have been transported from their original location by water erosion and deposited in a new sedimentary deposit. Clay deposits are typically associated with very low energy depositional environments such as large lakes and marine basins.

From prehistoric times, clay has been indispensable in architecture, in industry, and in agriculture. As a building material, it is used in the form of brick, either sun-dried or fired. Clays are also of great industrial importance, e.g., in the manufacture of tile for wall and floor coverings, porcelain, china, and earthenware, pipe for drainage and sewage. Highly absorbent bentonite is much used in foundry work for facing the moulds and preparing the moulding sands for casting metals. The less absorbent bentonite is used chiefly in the oil industry as filtering and deodorizing agents in the refining of petroleum. Clay is a good raw material for paper, ceramics, plastics, and rubber industries. Clays mixed with other materials serve as drilling muds to protect the cutting bit while drilling. Other uses are in the making of fillers, sizings, in purifying sewage. Clays are very useful raw materials from which many domestic and commercial wares can be manufactured (Adeola 2014, Oyinloye 1991, Bolarinwa and Adeola 2017). In Nigeria, clay deposits have not been utilized adequately considering the qualities of this type of industrial mineral that occurs in the country. This may be due to lack of geological information on the

assessment of the clay deposits and what they can be used to manufacture. It is very pertinent that the physical and chemical properties of any clay deposit should be ascertained for industrial uses and dressings in construction. The present investigation is intended to study the physical, chemical, mineralogical and industrial characteristics of the clay deposits in Papalanto and its environs appraise their economic potentials. This obviously will complement previous studies that were mainly on geochemistry of limestone, shale in the area.

2. Geological Setting

Papalanto and its environs is situated in the South-western part of Nigeria and falls within the Dahomey Basin. (Fig. 1)

The Dahomey Basin is a combination of inland/coastal/offshore basin that stretches from southeastern Ghana through Togo and the Republic of Benin to southwestern Nigeria. It is separated from the Niger Delta by a subsurface basement high referred to as Okitipupa Ridge. Its offshore extent is poorly defined. Sediment deposition follows an east-west trend. In the republic of Benin, the geology is fairly well known. In the onshore, cretaceous strata are about 20m thick. A non-fossiliferous basal sequence rests on the Precambrian basement. This is succeeded by coal cycles, clays and marls which contain fossiliferous horizons. Offshore, a 1000m thick sequence consisting of sandstones followed by black fossiliferous shale towards the top has been reported. This was dated by Billman (1976) as being pre-Albian to Maastrichtian. The cretaceous is divisible into two geographic zones, north and south. The sequence in the northern zone consists of a basal sand that progressively grades into clay beds with intercalations of lignite and shales. The uppermost beds of the Maastrichtian are almost entirely argillaceous. The southern zone has a more complicated stratigraphy with limestone and marl beds constituting the major facies. Sedimentation in the northern zone which is located inland and close to the basin periphery, began during the Maastrichtian when a thin sequence (<200m) of unconsolidated sands, grits, silts, clays and shales, was deposited. This sequence rests on the basement; the transitional facies is marked by a conglomerate or white to grey sandy and kaolinitic clays derived as degradation products from the surrounding Precambrian rocks. In the southern zone, which is coastal and offshore, the oldest sediments consist mainly of loose sand, grits, sandstones and clay with shale interbeds which progressively grade into shale. They are late Albian and possibly Neocomian in age (Omatsola and Adegoke, 1981). The basal conglomerates have been reported from outcrops and boreholes (Jones and Hockey, 1964). The onshore sequence towards the basin periphery in Nigeria correlates well with the Maastrichtian onshore in the republics of Benin and Togo. The geology of Togo sector is very similar to that of Nigerian and Benin sectors. The cretaceous succession shows marked lithological changes which have been expressed in terms of formal and informal lithostratigraphic nomenclature by previous workers. This can lead to dual or multiple nomenclature and thus confusion.

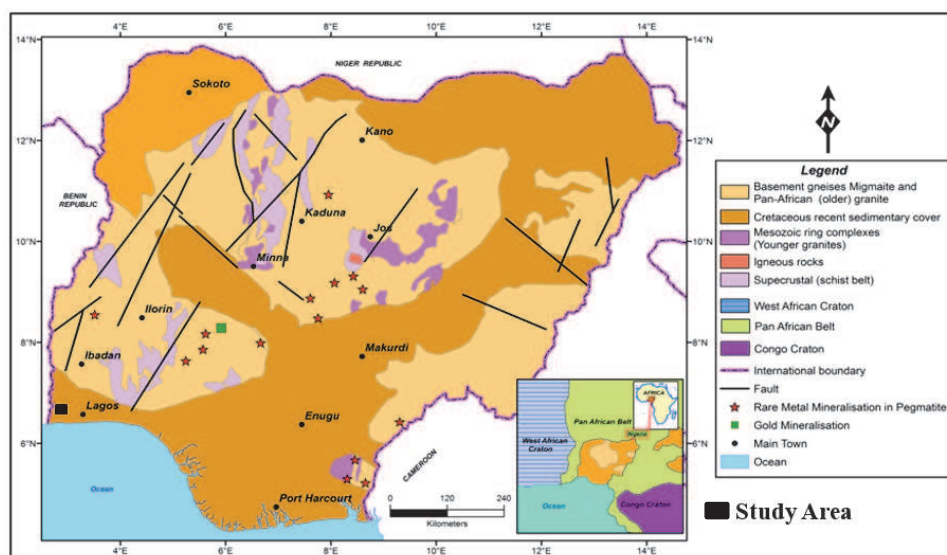


Figure 1. Map of Nigeria showing the location of the study area

Among the major lithostratigraphic units of the eastern Dahomey basin are the Araromi, Ewekoro, and Akinbo formations. The Dahomey basin is one of the sedimentary basins on the continental margin of the Gulf of Guinea, extending from southeastern Ghana in the west to the western flank of the Niger Delta (Jones and Hockey, 1964;

Omatsola and Adegoke, 1981. The basin is bounded in the west by faults and other tectonic structures associated with the landward extension of the fracture zone. Its eastern limit is similarly marked by the Hinge line, a major fault structure marking the western limit of Niger Delta (Adegoke, 1969; Omatsola & Adegoke, 1981). It is also bounded in the north by the Precambrian basement rock and the Bright of Benin in the south (Fig 2).

Stratigraphic studies of Dahomey basin were conducted by various researchers among whom are Jones and Hockey, (1964); Adegoke.(1975); Omatsola and Adegoke, (1981). The general sequence for the rock unit from the top are the Coastal plain sands, Ilaro formation, Oshosun formation, Akinbo formation, Ewekoro formation, and Abeokuta formation lying on the Southwestern Basement Complex of Nigeria (Fig. 3)

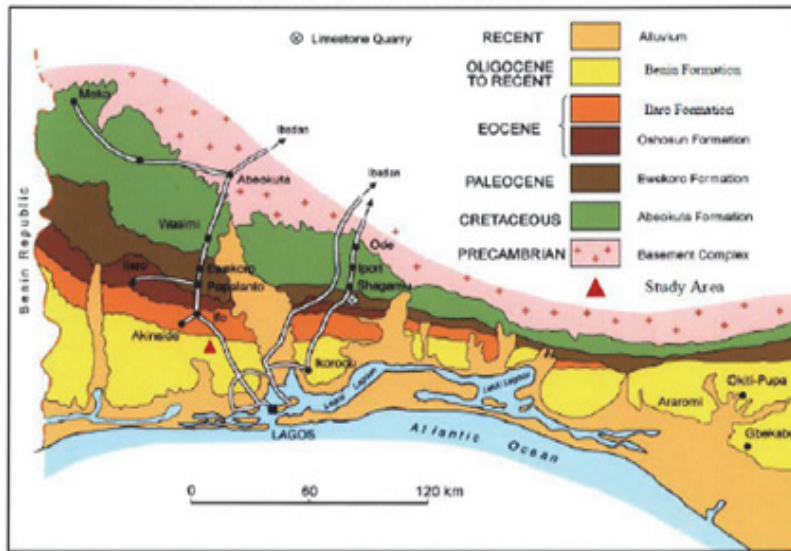


Figure 2. Map of Eastern Dahomey basin showing the stratigraphic setting

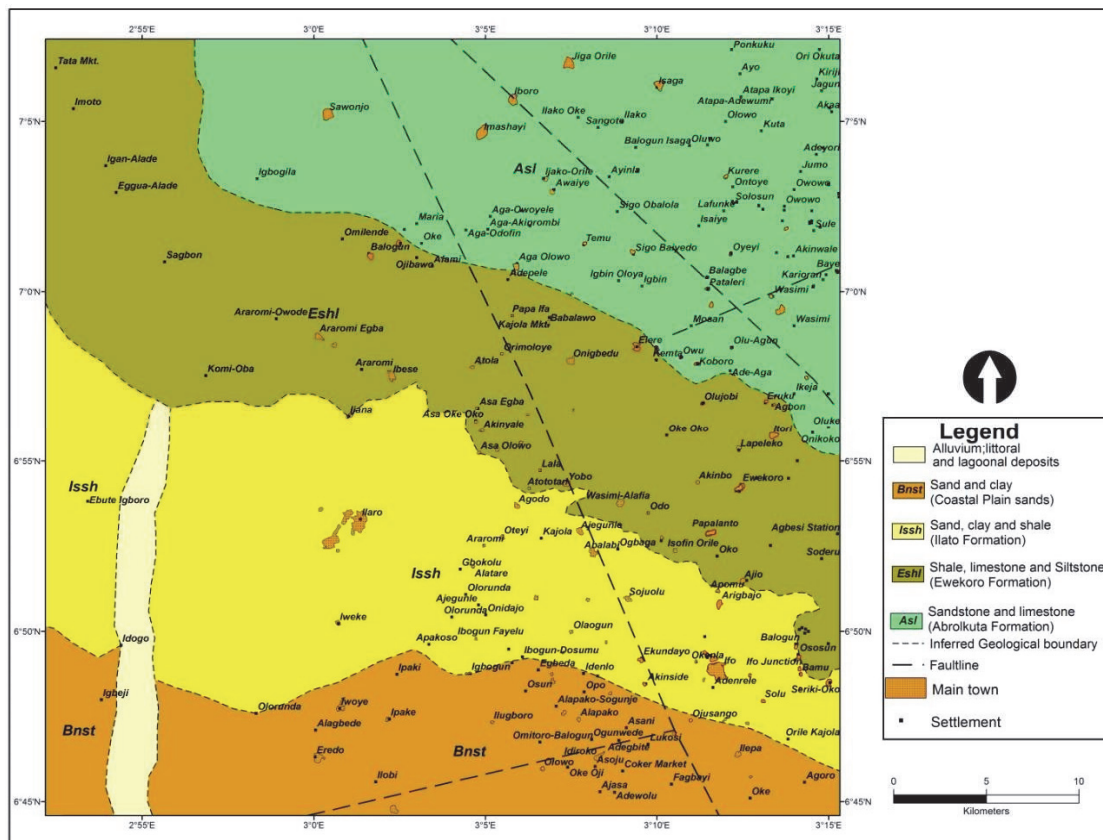


Figure 3. Geologic map of Papalanto and its environs

3. Materials and Method

Thirty two fresh representative samples were collected from the three locations. Eight samples from each of the locations were prepared for physical and chemical analyses. Hydrometer method and wet sieving analyses were used to determine the grain size distribution of Ifo, Papalanto and Imoto clays. Liquid and Plastic limits were determined using Cassangrade (1948) apparatus. Representative samples were made into pellets using a pressure gauge these pellets were fired at about 950°C for about 24 hours to determine the shrinkage capacity and colour change of the clay deposits. Samples were also selected for the thermal tests. Clay mineralogy was determined using X-ray Diffraction (XRD) at Activation Laboratory in Canada. The X-ray diffraction analysis was performed on a analytical X'Pert Pro diffractometer, equipped with a Cu X-ray source and an X'celerator detector, operating at the following conditions: voltage - 40 kV; current - 40 mA; range - 5-80 °2θ; step size: 0.017, °2θ; time per step: 50s; divergence slit: fixed, angle 0.5°. The crystalline mineral phases were identified by X'Pert HighScore Plus software using the PDF-4 ICDD database. The quantities of the crystalline minerals were determined using the Rietveld method. The Rietveld method is based on the calculation of the full diffraction pattern from crystal structure information. Elemental compositions of the clay samples were determined using Inductively Coupled Plasma-Mass Spectrometer (ICP-MS) at Activation Laboratory, Canada . For ICP-MS, microwave high pressure/temperature decomposition of samples (230°C, 7.0MPa; Paar Physical Multiwave sample preparation system) using Merck Suprapurs grade reagents (HF, HClO₄, HNO₃ and HCl). All measurements were made on a Sciex/Perkin-Elmer ELAR 6000 ICP-MS. The Chemical Index of Alteration (CIA = Al₂O₃ / Al₂O₃ + CaO + Na₂O + K₂O) X 100, where all components expressed in molecular proportions was calculated to measure the Intensity of Chemical Weathering (ICW) of soils.

4. Results and Discussion

4.1 Field work and Macro-Petrography

Profile 1 is located within the sedimentary basin of Nigeria at Ifo area, Southwestern Nigeria. It is located on latitude N06° 45' 75" and longitude E003° 12' 9". A 9m thick profile is exposed along Ifo-Abeokuta expressway. There is no evidence of large scale movement and the different horizons within the profile grades into each other. Three distinct layers were identified based on colour and texture. (Fig. 4a).

The upper horizon which is the topsoil is generally light brown in colour and is about 0.8m thick with texturally fine grained containing plant roots and humus. Underlying this horizon is a medium grained, reddish-brown laterite layer of about 4m in thickness. This layer contains no humus or organic materials like the topsoil. Underlying the laterite layer is the clay horizon. This clay is fine grained in texture and is whitish in color with some patches of purple and brown. It is about 5m in thickness.

Profile 2 is located at Papalanto area on latitude N06° 53' 28" and longitude E003° 07' 84".

A 8.5m thick profile is exposed at Ajegunle Bus-Stop along Papalanto-Ilaro road, There is no evidence of large scale movement and the different horizons within the profile also grades smoothly into each other like Ifo profile. Based on colour and textural characteristics, three distinct layers were identified (Fig. 4b).

The upper horizon which is the topsoil is about 0.5m thick and brown in colour. It is fine grained in texture and contains lots of plant roots and humus. Directly below this horizon is the reddish-brown laterite layer of about 5m in thickness and grades gradually into the underlying clayey horizon. The clay deposit is vast in this area. This clay deposit is fine grained in texture and whitish in colour with some patches of colour purple and brown.

Profile 3 is located within the sedimentary terrain at Imoto area. It is located on latitude N07° 05' 43.6" and longitude E04° 18' 9.52". A 5m thick profile is exposed in a quarry near Igan-Alade, Igbogila. Based on colour and textural characteristics, two distinct layers were identified (Fig. 4c). The topsoil is light brown in colour and is about 0.5m thick. It is fine grained in texture and contains lots of plant roots and humus. Directly below this horizon is the clayey horizon. This clay is fine grained in texture and is dark grey in colour.

4.2 Mineralogy

The X-ray diffractograms derived from Ifo, Papalanto and Imoto clays show that kaolinite is the dominant clay mineral in the profiles. Conspicuous peaks of kaolinite are recorded at $d = \sim 7.14 \text{ \AA}$, $d = \sim 4.36 \text{ \AA}$ and $d = \sim 3.57 \text{ \AA}$ in the oriented and non-glycolated samples. Quartz is the major non clay mineral that is present in all the horizons in the profiles because of its high resistance to weathering. Anatase and hematite occur as traces in all the profiles in the area (Fig 5).

4.3 Chemical Compositions

The average chemical compositions of major and trace elements in the clays are presented in Table 1. The mean

compositions of SiO₂ in Ifo, Papalanto and Imoto area are 60.35%, 56.47%, and 61.95% respectively. This shows that the SiO₂ in Imoto clay profile is very high. This might be due the presence of high concentration of quartz as recorded in Table 1. The enhanced value of 61.95% in the clay may be due to relative depletion of MnO, MgO, CaO, Na₂O, K₂O in the horizon and could be said that quartz which is of secondary origin probably accumulated from the chemical weathering of rock forming silicates. The weathering and dissolution of silicate minerals consequently led to the enrichment of SiO₂ and Fe₂O₃ in the topsoil. The X-ray diffractogram of the clay show prominent peaks of quartz probably due to their relative crystallinity (Brindly 1961).

The concentration of Al₂O₃ ranges from 15.57 to 26.15%. The aluminium concentration in Papalanto clay is the highest (26.15%) and the lowest values are recorded in Imoto clay (15.57%). Fe₂O₃ is generally low in the clays in this area except in Imoto clay. The concentration of Fe ranges from 2.17 to 6.72%. The brownish colour of moto clay can be attributed to the presence of high concentration of iron as evidenced in Table 1. This is strongly supported by several peaks of hematite in the X ray diffractogram of Imoto clay (Fig. 5).

The following low values CaO, Na₂O, K₂O, P₂O₅, MgO and LOI strongly indicate leaching during chemical weathering. The concentration values for CaO are 0.02%, 0.02%, and 0.17%, Na₂O are 0.02%, 0.04%, and 0.10%, and P₂O₅ are 0.07%, 0.06% and 0.05% in Ifo, Papalanto and Imoto respectively. TiO₂ is relatively low in all the clays but high in Imoto clay with average of (1.96%). P₂O₅ is relatively low in all the clays.

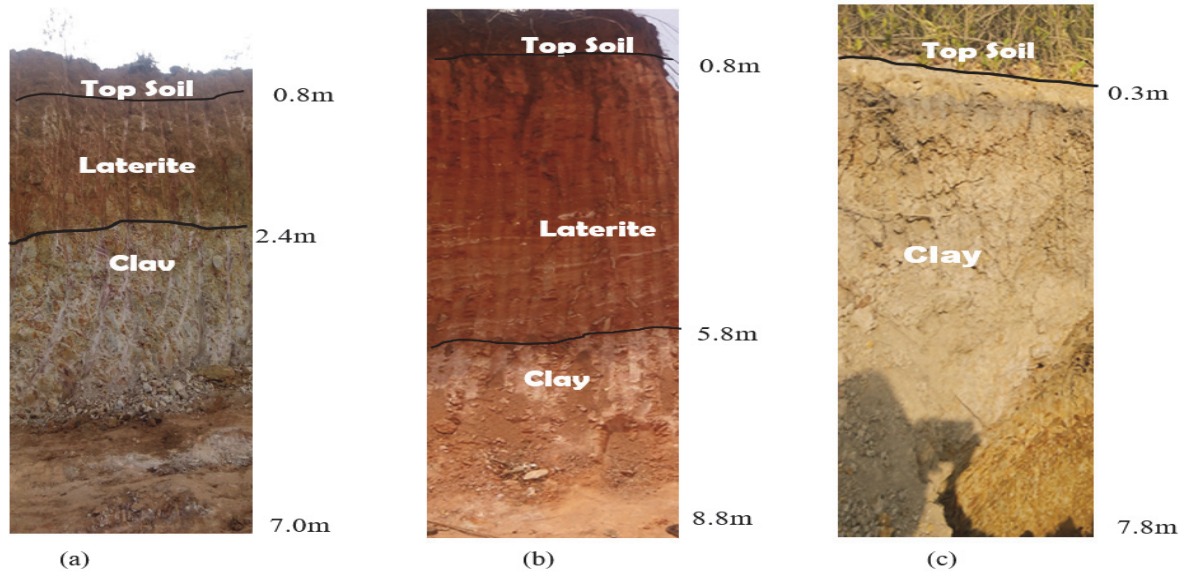
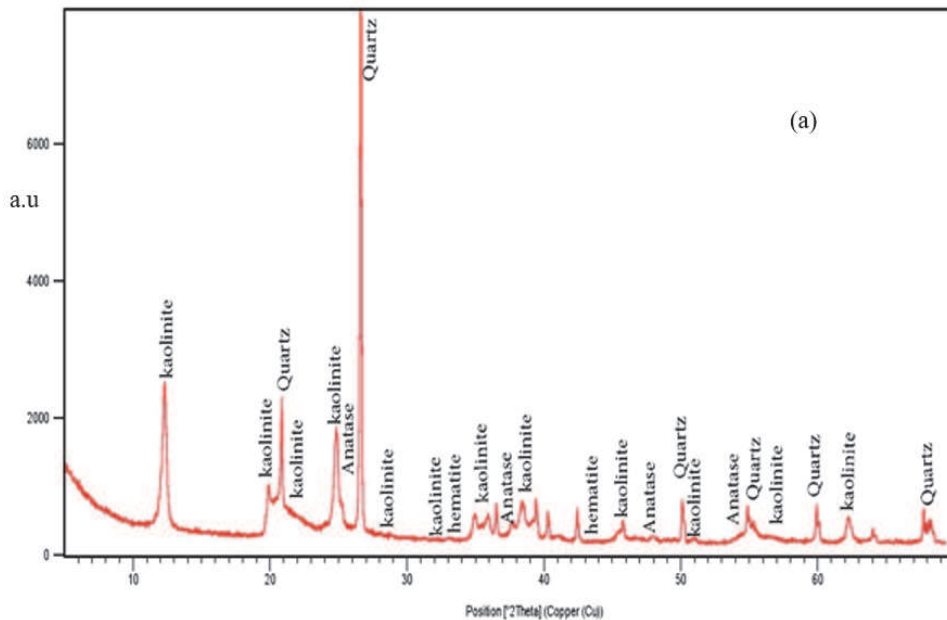


Figure 4. Weathered profiles of the study area. (a) - Ifo clay body, (b) - Papalanto clay, (c) - Imoto clay body



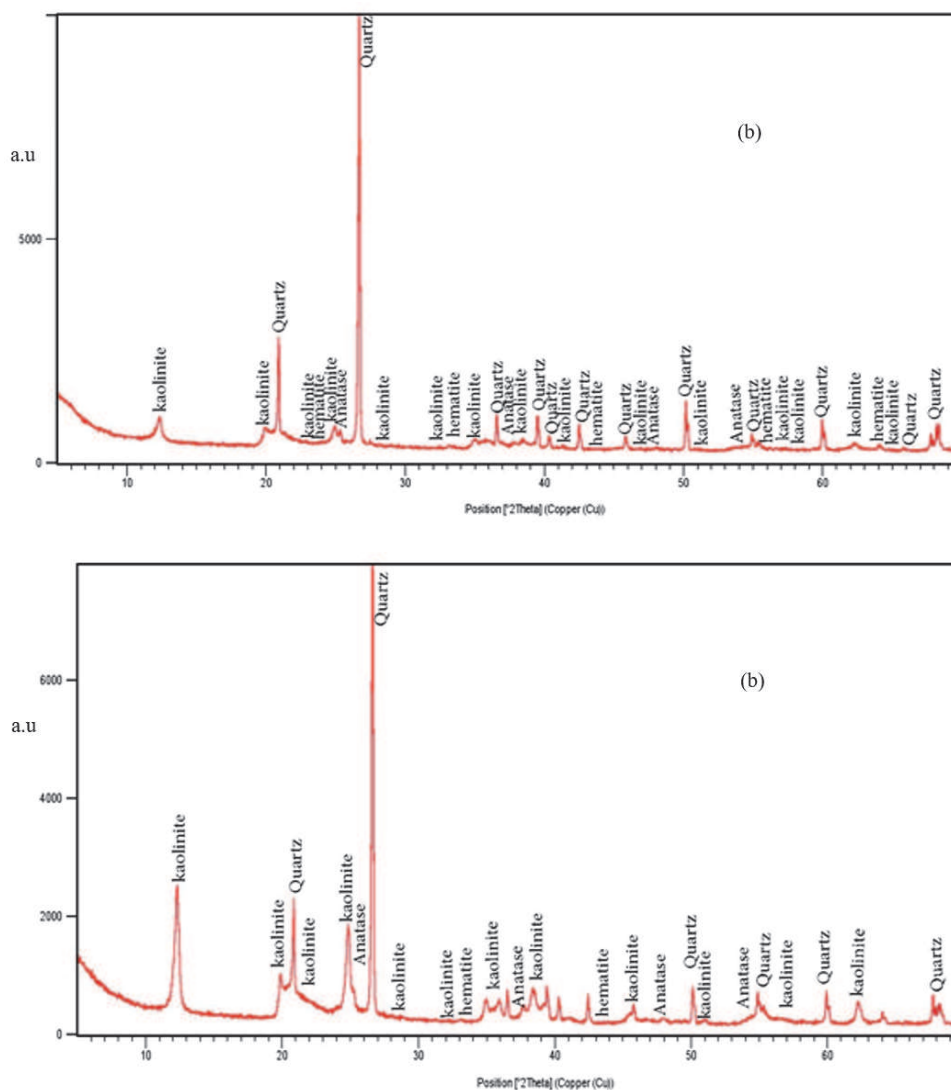


Figure 5. X-ray diffraction diagrams for the clay samples (a) Ifo clay, (b) Papalanto clay, (c) Imoto clay

The trace element Ba is relatively high in Ifo clay (129ppm) and Papalanto clay (123ppm) but low in Imoto clay with average of (68.60%). Zr and V are relatively high while Sc, Be showed low values in Ifo, Papalanto and Imoto clays. Silica ratio (S.R) is high in Imoto clay but relatively low in Ifo clay and Papalanto clay with averages of (4.61%) and (4.85%) respectively. Alumina ratio (A.R) is relatively low in Imoto clay with average of (2.31%) but high in Ifo clay (11.41%) and Papalanto clay (9.72%).

Major elements abundances show that the clay samples have SiO₂ (ca. 59.46%), Al₂O₃, (ca 22.16%), and Fe₂O₃ (ca 3.06%) constituting 98.3% of the bulk compositions. Al₂O₃ content is higher in Ifo and Papalanto but lower in Imoto. The Fe₂O₃ is relatively higher in Imoto compared with other clays. Fe₂O₃ MgO, MnO, CaO Na₂O and TiO₂ are higher in Imoto clay than clay samples from Papalanto and Ifo.

The SiO₂, Al₂O₃ and Fe₂O₃ of Papalanto clay compared favourably with Plastic Fire clay (Huber, 1985) as presented in the chemical compositions (Table 2) except that the Al₂O₃ is slightly higher. In terms of functional applications, Ifo Papalanto and Imoto chemical compositions are within the limits of industrial applications for ceramics and refractories (Table 3).

The Chemical Index of Alteration (CIA = Al₂O₃ / Al₂O₃ + CaO + Na₂O + K₂O) X 100, where all components expressed in molecular proportions is commonly used as a measure of the Intensity of Chemical Weathering (ICW) of soils. This parameter was derived by (Nesbitt and Young 1982, 1984). Nesbitt indicated that CIA values between 30-55 is an indication of weathering at incipient zone, and CIA values ranging between 51-85

could be considered as intermediate zone of weathering while weathering that is at the advanced stage will have CIA values greater than 85. From the values of Chemical Index of Alteration (CIA) calculated above, the CIA values for Ifo, Papalanto and Imoto clays are 99.27%, 98.75% and 95.28% respectively. It can be inferred that the weathering that resulted in the formation of these clays have reached advanced stage because all the CIA values are greater than 85% (Table 1).

4.4 Industrial Properties

The industrial properties of clay such as thermal characteristics, loss on ignition, fired colour, water absorption capacities and shrinkage were evaluated. The summary of the results of the industrial properties are presented in table 4. Loss on ignition ranges between 11.65%- 11.96% with a mean of 11.75%. The average values of water absorption capacities of fired pellets after immersion in cold water for 24 hours range from 12.09% – 17.14% with Imoto (17.14%) having the highest water absorption capacity.

The concentration of the iron compounds that are present in the samples determined the colour of the fired samples that ranged from brown to reddish brown. The average linear shrinkage is higher in Ifo and Papalanto (15%) clay samples than Imoto Clay (11%). The results of the wet analysis show that the average clay fraction of (Papalanto (85.4%) is greater than Ifo (69.4%) and Imoto (44.5%) clay. Plots of the plasticity indices against liquid limits values (Fig. 7) classify Papalanto, Ifo and Imoto clays as inorganic clays and inorganic silts with moderate compressibility and moderate toughness.

Table 1. Average chemical compositions (%) of major and minor oxides of clays in study area

Oxides	Ifo		Papalanto		Imoto	
	Mean	Range n=3	Mean	Range n=3	Mean	Range n=3
SiO ₂	60.35	59.57-61.97	56.47	53.71-57.77	61.95	61.52-62.52
Al ₂ O ₃	24.76	23.61-25.47	26.15	25.40-27.13	15.57	15.32-15.80
Fe ₂ O ₃	2.17	2.14-2.19	2.69	2.41-3.25	6.72	6.42-6.96
MnO	0.01	0.01-0.01	0.02	0.02-0.02	0.03	0.03-0.03
MgO	0.05	0.04-0.05	0.06	0.06-0.06	0.41	0.41-0.42
CaO	0.02	0.02-0.02	0.02	0.02-0.02	0.17	0.17-0.17
Na ₂ O	0.02	0.02-0.02	0.04	0.03-0.04	0.10	0.10-0.10
K ₂ O	0.09	0.09-0.01	0.21	0.19-0.23	0.09	0.09-0.10
TiO ₂	1.73	1.39-1.88	1.58	1.54-1.65	1.96	1.94-1.98
P ₂ O ₅	0.07	0.05-0.10	0.06	0.05-0.06	0.05	0.05-0.06
LOI	10.55	10.38-10.64	11.93	11.59-12.47	11.99	11.80-12.33
TOTAL	99.82	99.31-100.3	99.23	98.50-100.5	99.06	98.40-99.40
Trace Elements (ppm)						
Ba	129.00	99-140	123.00	114-128	68.60	68-70
Sr	104.30	56-123	49.00	47-50	46.60	46-48
Y	25.50	20-28	30.00	29-32	58.00	58-58
Sc	11.50	10 --12	12.00	12--12	19.00	19-19
Zr	920.50	698--1042	968.00	949-978	945.60	935-958
Be	1.00	1--1	1.00	1--1	4.00	4--4
V	129.00	97-140	104.00	98-115	130.60	128-132
Silica and Alumina Ratio %						
SR	4.61	4.52-4.79	4.85	4.65-5.22	10.69	10.37-10.90
AR	11.41	10.88-11.68	9.72	8.34-10.56	2.31	2.24-2.46
MgO + CaO	0.07	0.06-0.07	0.08	0.08-0.08	0.58	0.58-0.59
Na ₂ O+K ₂ O	0.11	0.11-0.12	0.20	0.5-0.27	0.19	0.2-0.19
CIA	99.27		98.75		95.28	

Table 2. Average chemical composition of Ifo, Papalanto and Imoto residual clays

Oxides	*Ifo (%)	*Papalanto (%)	*Imoto (%)	Reference (A) %	Reference (B) %	Reference (C) %
SiO ₂	60.35	56.47	61.95	52.92	57.67	46.88
Al ₂ O ₃	24.76	26.15	15.57	9.42	24.00	37.65
Fe ₂ O ₃	2.17	2.69	6.72	3.65	3.23	0.88
MgO	0.05	0.06	0.41	0.08	0.30	0.13
CaO	0.02	0.02	0.17	1.91	0.70	0.03
Na ₂ O	0.02	0.04	0.10	0.03	0.20	0.21
K ₂ O	0.09	0.21	0.09	0.98	0.50	1.60
P ₂ O ₅	0.07	0.06	0.05	0.02	-	-
MnO	0.01	0.01	0.03	-	-	-
LOI	10.47	11.93	11.99	10.19	10.50	12.45
Total	98.01	97.64	97.03	79.20	97.10	99.83

*Average values for 5 samples

(A) - Florida Active Kaolinite (Huber, 1985)

(B) - Plastic Fire Clay, St Louis (Huber, 1985)

(C) - China Clay (Huber, 1985)

Table 3. Comparison of Ifo, Papalanto and Imoto clays with some industrial chemical specifications

Oxides	*Ifo (%)	*Papalanto (%)	*Imoto (%)	Reference		
				(A) %	(B) %	(C) %
SiO ₂	60.35	56.47	61.95	47.90-48.30	67.57	51.0-70.0
Al ₂ O ₃	24.76	26.15	15.57	37.90-38.40	26.50	25.0-44.0
Fe ₂ O ₃	2.17	2.69	6.72	13.40-13.80	0.50-1.20	0.2-0.7
MgO	0.05	0.06	0.41	0.20-0.30	0.10-0.19	0.2-0.7
CaO	0.02	0.02	0.17	0.03-0.25	0.18-0.30	0.1-0.2
Na ₂ O	0.02	0.04	0.10	0.20-0.35	0.20-1.50	0.8-3.5
K ₂ O	0.09	0.21	0.09	0.40-0.10	1.10-3.10	-
P ₂ O ₅	0.07	0.06	0.05	0.02	-	-
MnO	0.01	0.01	0.03	-	-	-
LOI	10.47	11.93	11.99	-	-	-
TOTAL	98.01	97.64	97.03	-	-	-

*Average values for 5 samples

(A) - Paints (Payne, 1961)

(B) - Ceramics (Singer and Sonja, 1971)

(C) - Refractory Bricks, (Parker, 1967)

Table 4. Physical and firing properties of the Ifo, Papalanto and Imoto clays

Profile	Clay Fraction	Atterberg Limits			LOI (%)	LSK(%)	WAC(%)	S.G	Colour
		LL (%)	PL (%)	PI (%)					
* Ifo clay	69.4	48.96	28.55	19.96	11.65	15	14.29	2.57	Brown
*Papalanto	85.4	52.48	34.04	20.14	11.65	15	12.09	2.55	Buff brown
*Imoto	44.5	49.16	28.72	21.15	11.99	11	17.14	2.60	Brown

* Average values for 5 samples

LOI – Loss on ignition

LSK – Linear Shrinkage

WAC – Water absorption capacity

LL – Liquid limit

PL – Plasticity limit

PI – Plasticity index

5. Discussion

Papalanto, Ifo and Imoto clays are basically residual and dominantly kaolinitic. Ifo and Papalanto clays are the most plastic while Imoto is sandy and characterized with low plasticity. The presence of hematite indicates that the ferromagnesian mineral such as biotite has been completely weathered and converted to hematite. This is actually responsible for the brown colouration in the clays.

Comparing the chemical composition with the standard industrial references, Ifo, Papalanto and Imoto clays cannot be used for the production of paint because of the high silica and low alumina concentration but they are quite recommended for the production of ceramics. Also, with further beneficiation, all of them can serve as good raw materials for the production of refractory bricks.

The composition of the Papalanto clay is quite comparable with Plastic fire clay (Huber, 1985) with further beneficiation and Ifo and Imoto clays can meet the Florida active kaolinite standard. None of the clays can be used as China clays because of the high silica and alumina content.

In determining the accurate engineering properties of soil, the proposed Cassangrade (1948) plasticity chart is generally employed (Fig.7). The chart incorporates a boundary called “A-Line” which starts from the liquid limit of 20 and runs diagonally upward to the right with a slope of 0.73. The line demarcates the inorganic clays from inorganic silts and organic clays. From the result, the liquid limit is higher than the plastic limit (Table 4). This agrees with the general trend observed by Adeyemi (2001). The difference between these two limits called the Plasticity Index is an important geotechnical parameter. When the value of the plastic limit is close to that of the liquid limit, soil is said to be non-plastic. From this result, none of the plasticity and plasticity index values is close to each other. Consequently, the clays in Papalanto area can be said to be plastic. Plots of plastic limit and plasticity indices on the Bain (1971) mouldability chart (Fig. 7) indicate that the clay bodies from Pakkshin area generally possess moderate mouldability. It must be noted that the shrinkage and mouldability of the clays, which largely determined the industrial suitability, are affected by the grain size distribution and the mineralogy of the clay. Kaolinitic clays have no expandable lattice and therefore have low swelling potential and low shrinkage. Clayey materials in this area are mainly kaolinite. This is similar to the mouldability of clays above biotite granite, banded gneiss in Abeokuta area by Bolarinwa (2001).

Based on the percentage of clay, silt, and sand, the textural classification of clays in the study area, Ifo clays plot within the field of (Sandy silt) zone. Papalanto and Imoto clays plot within the field of Sandy mud and sandy silt (Fig.6) due to their moderately high silt and mud. The grain size distribution curves obtained showed that the soils are well graded. The particle size distribution is evenly spread between sand, silt and clay. High amount of coarse fractions in a soil sample and low amount of fines are generally believed to give good subgrade soils as they contribute significantly to the mechanical strength of the soil. These clays would not serve as good subgrade material although due to their high specific gravity but they can be used as filling materials.

Fabbri and Fiori (1985) constructed ternary diagrams based on the chemical and mineralogical compositions of clay raw materials to be used in the ceramics industry for the production of traditional ceramic products. These ternary diagrams are the most suitable to define and visualize compositional fields. Based on this, only Imoto clay falls near the global fields of red-stone ware and this suggests that Imoto clay is the only suitable material for the production of red-stone ware products. The chemical composition of Imoto is quite similar to the clay materials from Bailen area and Greece being used for the production of Red-stoneware as reported by Gonzalez et al. (1997) and Oikonomopoulos et al. (2007). The amount of alkalis ($\text{Na}_2\text{O} + \text{K}_2\text{O}$) concentration in Imoto, Papalanto and Ifo clays is favourably compared with Moro and Chulilla clays employed for the production of ceramic in Castellon (Mesenguer 2011).

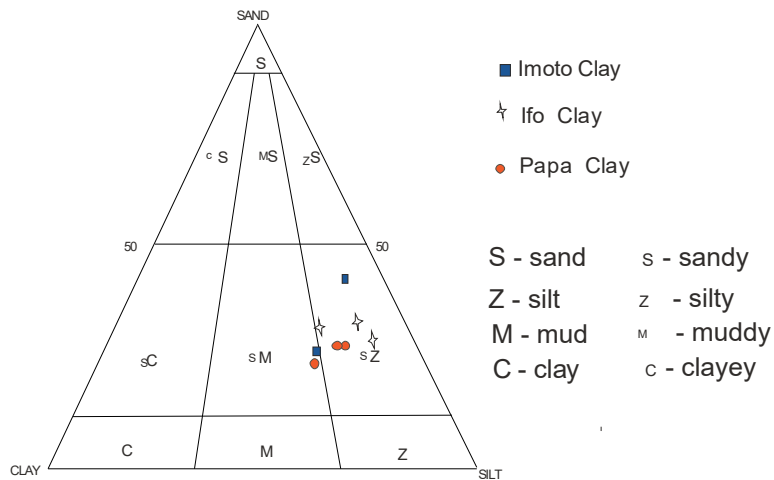


Figure 6. Textural classification of Ifo, Papalanto and Imoto clays (After Folk, 1974)

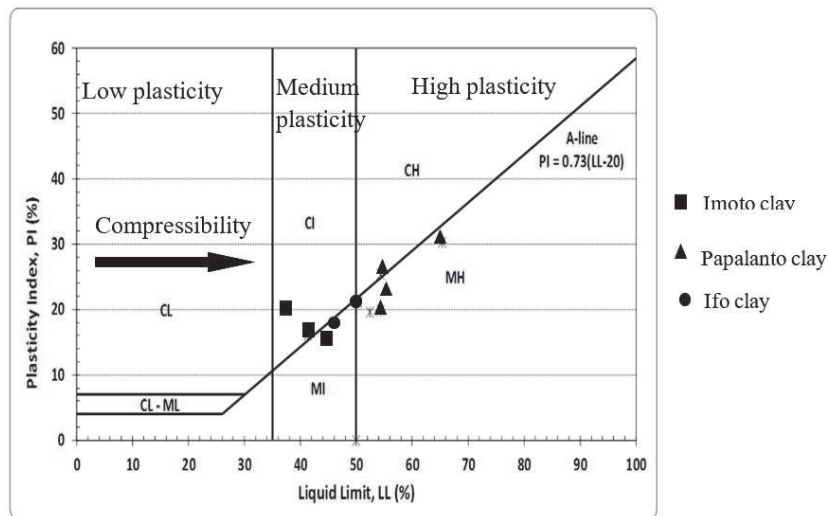
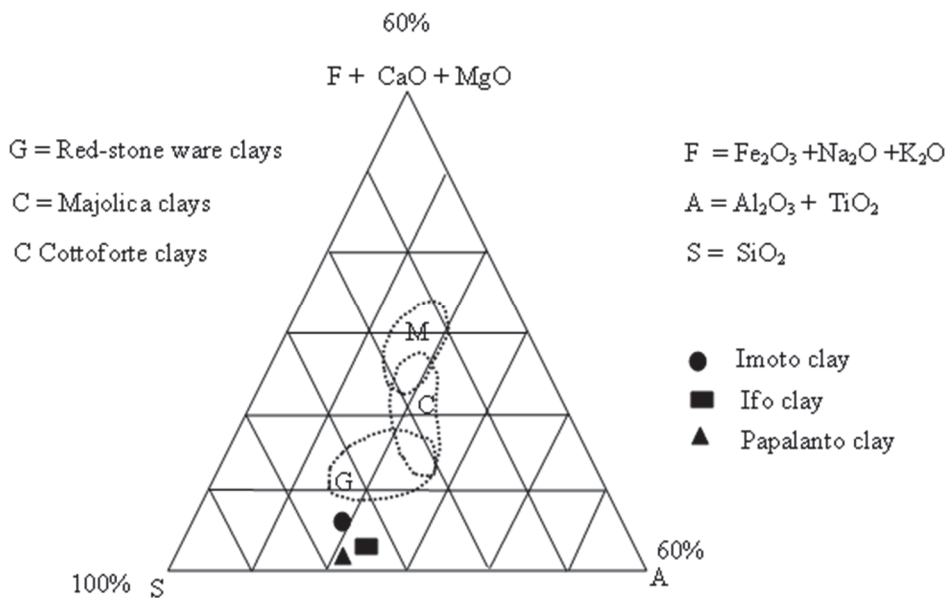


Figure 7. A plasticity chart for the classification of clays from Ifo, Papalanto and Imoto area (after Cassangrande, 1948)



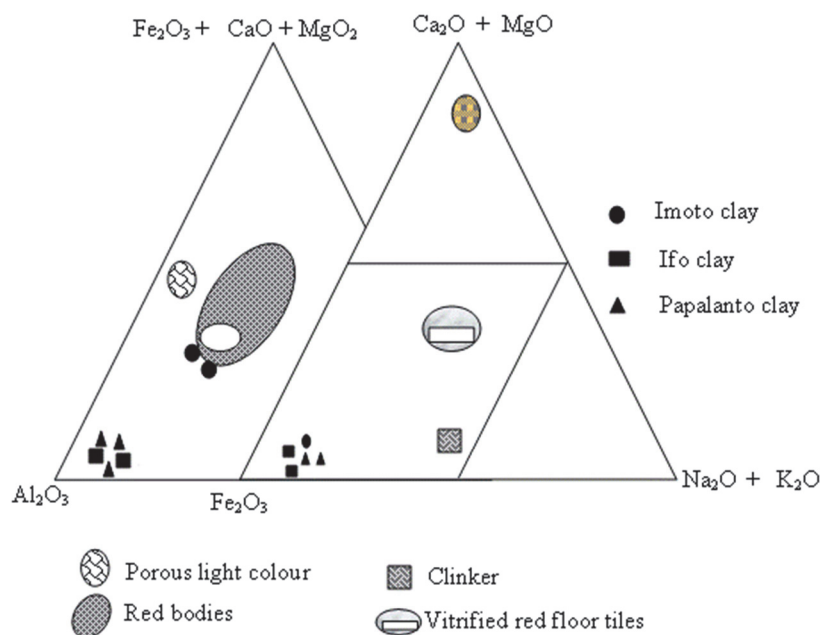


Figure 8. Ternary diagram $\text{SiO}_2/\text{Al}_2\text{O}_3 + \text{TiO}_2/\text{Fe}_2\text{O}_3 + \text{MgO} + \text{K}_2\text{O} + \text{CaO} + \text{Na}_2\text{O}$ (Afer Fabbri and Fiori, 1985)

6. Conclusion

The mineralogical study in this area shows that the kaolinite is the dominant mineral in the clay and other non-clay minerals like quartz, hematite and plagioclase occurs as accessory minerals. From the mineralogical and chemical point of view, Imoto favours the production of ceramics and paplanto, and Ifo would be adequate for the production of refractories and burnt bricks. Open pit method is hereby recommended for the mining of these clay deposits because they are generally covered with thin soil layer ranging from 0.3m to 0.8m in thickness.

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